

El paisaje de las montañas atlánticas: avances en el conocimiento geohistórico y ambiental



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Palaeoecological review of the climatic, cultural and plant landscape dynamics of the Cantabrian region

Revisión paleoecológica de la dinámica climática, cultural y del paisaje vegetal de la región cantábrica

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Resumen: La región del noroeste peninsular Ibérico ha demostrado ser un excelente laboratorio a campo abierto para el estudio de las dinámicas del paisaje vegetal a lo largo del tiempo. El gran número de cuevas, lagos y turberas, entre otros yacimientos existentes en la región, ha permitido el análisis de múltiples registros sedimentarios y arqueológicos a partir de técnicas palinológicas, análisis de carbones sedimentarios y de otros indicadores paleoambientales. En este sentido, durante las últimas décadas se han aportado distintas evidencias sobre cómo han evolucionado las comunidades vegetales durante el Último Periodo Glacial y el Holoceno. Además, también se han estudiado factores clave implicados en la configuración del paisaje, tales como el clima, los incendios o las distintas fases culturales que han tenido lugar en la región. Con el fin de contextualizar en qué punto se encuentra el conocimiento paleoecológico del noroeste peninsular Ibérico y más en concreto el de la región de Cantabria, el presente trabajo revisa el estado de la cuestión de tres grandes temáticas paleoambientales: la

evolución del clima, las etapas culturales del ser humano y el estudio de las dinámicas del paisaje vegetal a partir de registros paleoecológicos.

Palabras clave: Península Ibérica, Cantabria, Último Periodo Glacial, Holoceno, Dinámicas del paisaje vegetal, Paleoecología, Clima, Influencia antrópica, Geografía histórica ambiental.

Abstract: The northwestern Iberian Peninsula has proved to be an excellent open field laboratory for the study of plant landscape dynamics over time. The large number of caves, lakes and peat bogs, among other sites in the region, has allowed the analyses of multiple sedimentary and archaeological records using palynological techniques, sedimentary charcoal analyses and other palaeoenvironmental indicators. In this sense, over the last few decades, different evidence has been provided on how plant communities evolved during the Last Glacial Period and the Holocene. In addition, key factors involved in the configuration of the landscape have also been studied, such as climate, fires or the different cultural phases that have taken place in the region. In order to contextualise the current state of palaeoecological knowledge of the northwestern Iberian Peninsula and more specifically of the Cantabrian region, this work reviews the state of the art of three major palaeoenvironmental topics: the evolution of climate, the cultural stages of human beings and the study of the plant landscape dynamics based on palaeoecological records.

Key words: Iberian Peninsula, Cantabria, Last Glacial Period, Holocene, Plant landscape dynamics, Palaeoecology, Climate, Anthropogenic influence, Historical environmental geography.

PALAEOCLIMATOLOGY IN THE NORTH ATLANTIC REGION AND AN APPROACH TO THE CANTABRIAN RANGE

Glacial-Interglacial cycles and the Last Glacial Period

Climate during the Quaternary has been characterised by an alternation between cold and warm intervals. Such oscillations have been studied through several methods in fossil cores, and oxygen isotopes are one of the proxies most commonly used.

For ocean sediments, Lisiecki & Raymo (2005) presented a 5.3-Myr stack of 57 $\delta^{18}\text{O}$ records for all over the world (Fig. 1). The sequences were

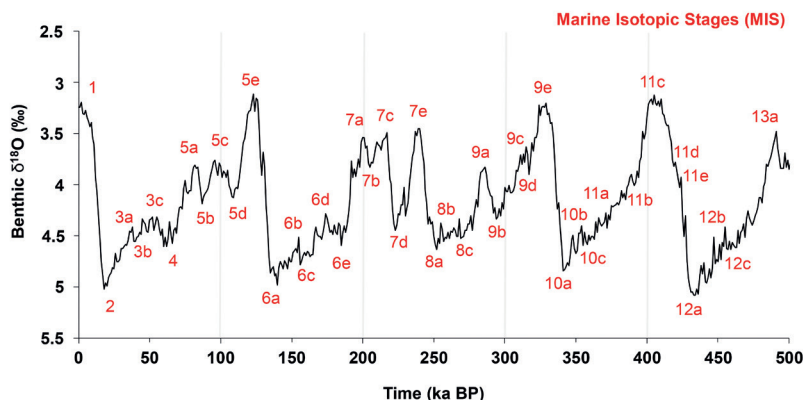


FIGURE 1. Benthic $\delta^{18}\text{O}$ from the LR 04 stack for the 500 ka BP. In red: the onset of Marine Isotopic Stages (MIS) and substages identified within them. Adapted from: Lisiecki & Raymo (2005) and Railsback et al. (2015).

globally well-distributed and covered the entire Quaternary period (Holocene and Pleistocene epochs) and the Pliocene epoch. In total, 104 Marine Isotopic Stages (MIS) were identified along the Quaternary. Some MIS were also subdivided into lettered substages considering fluctuations from other environmental proxies (e.g., pollen and loess), from both oceanic and non-oceanic records (Railsback et al., 2015). According to that classification, the Holocene corresponds to MIS-1 (c. 0 – 11.7 ka BP) and the Last Glacial Period, also known as the Würm Glaciation, encompasses the MIS-2 (c. 11.7 – 28 ka BP), MIS-3 (c. 28 – 60 ka BP), MIS-4 (c. 60 – 75 ka BP) and MIS-5a-d (c. 75 – 117 ka BP). MIS-5e (c. 117 – 130 ka BP), which is the oldest substage of MIS-5, corresponds to the Eemian interglacial.

Ice records have also registered detailed climate variability from the present to the Eemian interglacial. A climate classification was established from the study of oxygen isotopes into different periods, known as the Dansgaard–Oeschger (DO) events (Fig. 2; Dansgaard et al., 1982; Johnsen et al., 1992). These episodes were divided into Stadials (cold conditions) and Interstadials (warm conditions), and as well as marine isotopic stages, some of them were also subdivided. More specifically, the DO events detected in Greenland ice cores are known as Greenland Stadials (GS) and Greenland Interstadials (GI). A common nomenclature was established by the synchronisation of the GRIP, GISP2 and NGRIP Greenland ice cores (Fig. 2; Rasmussen et al., 2014; Seierstad et al., 2014). In total, 26 GS and 25 GI, with their corresponding subdivisions, were identified along the Last Glacial Period. Additionally, occasional events of massive ice-rafted debris (IRD) occurred during some Stadials as the result of North Atlantic glaciers melting, which are known as Heinrich events or Heinrich layers (Fig. 2; Heinrich, 1988; Bond et al., 1992; Hemming, 2004).

Within the Last Glacial Period, most ice sheets reached their maximum extent between c. 26.5 and 19 ka BP, a cold period which is known as the Last Glacial Maximum (Clark et al., 2009). The harsh conditions led to the extension of the ice and permafrost cover in northern Europe (Clark et al., 2002; Vandenberghe et al., 2014; Lindgren et al., 2016). In the Cantabrian range a glacier advance has also been detected between c. 22.5 to 18 ka BP, even though the greatest extension of glaciers (Local Last Glacial Maximum) occurred during the initial phases of the Würm glaciation (Frochoso et al., 2013; Ruiz-Fernández et al., 2016). At all events, the sea level reached its maximum depth throughout the Last Glacial Maximum, approximately 120 – 135 m deeper than nowadays (Lambeck et al., 2014). At a global scale, a stack of the sea level variations from the last 798 ka was reconstructed considering 12 ocean records (Fig. 2; Spratt & Lisiecki, 2016).

Environmental conditions became warmer after the Last Glacial Maximum. The Bølling-Allerød chronozone, which corresponds to the DO-1, was a warm period between 14.6 – 12.9 ka BP (Iversen, 1953). Within this episode, seven subevents were established within the GI-1, indicating climate variability over a brief period of time (Rasmussen et al., 2014). Between the Bølling-Allerød chronozone and the onset of the Holocene lies the last Stadal of the Last Glacial Period, the Younger Dryas (Iversen, 1953). It was a cold period triggered by the discharge of meltwater into the oceans (Fairbanks, 1989; Teller & Leverington, 2004), affecting the North Atlantic thermohaline circulation and subsequently

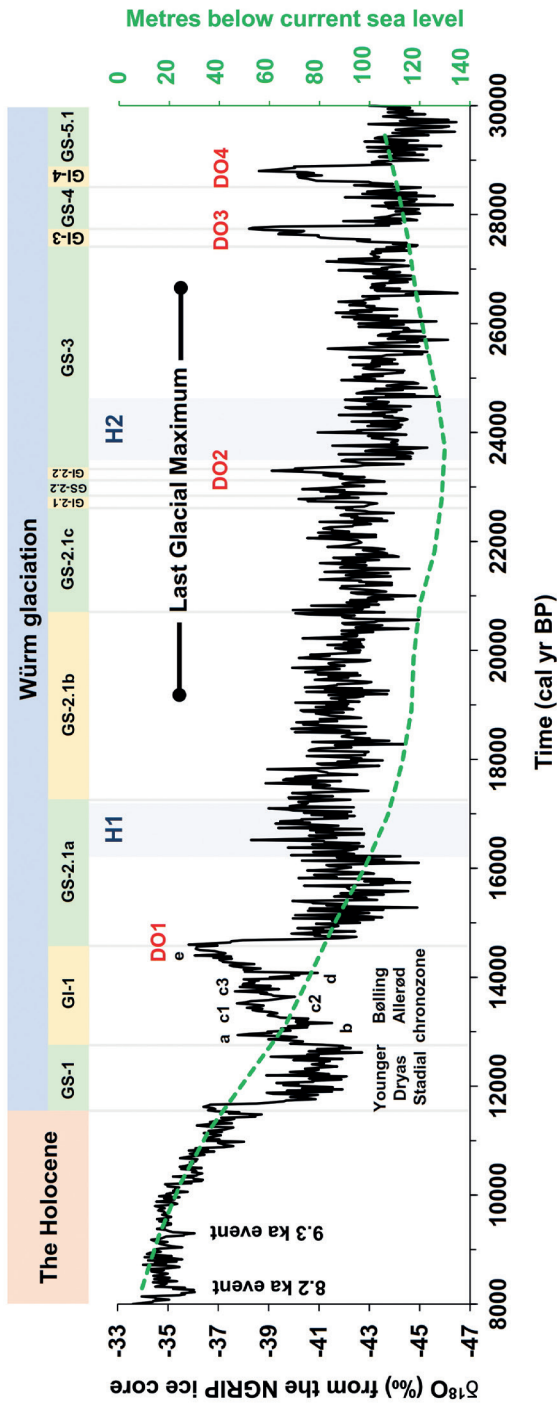


FIGURE 2. Climate proxies from 30,000 to 8000 yr BP. In a black line: $\delta^{18}\text{O}$ from the NGRIP core, adapted from Rasmussen et al. (2014). At the top of the figure, the chronology based on Greenland Stadials (GS) and Greenland Interstadials (GI) is shown, adapted from Rasmussen et al. (2014). In a green dashed line: metres below the current sea level, modified from the sea level stack obtained in Spratt & Lisiecki (2016). In blue rectangles: Heinrich events identified and discussed by Heinrich (1988); Bond et al. (1992) and Hemming (2004). In red: Dansgaard – Oeschger events (DO) identified by Dansgaard et al. (1982) and Johnsen et al. (1992).

the atmospheric climate (Clark et al., 2001, 2002). Several studies have examined the impact that such climate variability had on the Cantabrian region through isotopic analyses of cave speleothems (Domínguez-Villar et al., 2008, 2009, 2017; Moreno et al., 2010a; Baldini et al., 2015, 2019; Smith et al., 2016; Rossi et al., 2018). The transition from the Bølling-Allerød chronozone to the Younger Dryas Stadial entailed a reduction in temperature ranging from 6 to 9 °C on the Cantabrian littoral, a significant temperature relapse that was accompanied by winter aridity (Figs. 2 and 3; Baldini et al., 2015, 2019).

The Holocene

The Holocene is the Interstadial that runs until today. The onset of this epoch is dated to 11.7 ka BP (Walker et al., 2009), when a strong climate shift occurred in a brief period of time. Climatologically, this event brought about alterations in atmospheric circulation (Steffensen et al., 2008) and air temperature, which was estimated to increase by about 10 ± 4 °C (Severinghaus et al., 1998; Grachev & Severinghaus, 2005) in less than 50 years (Taylor et al., 1997; Alley, 2000). As a result, several environmental proxies such as the $^{18}\text{O}/^{16}\text{O}$ ratio, ice chemistry and the dust accumulation rate registered sharp variations (Johnsen et al., 2001; Steffensen et al., 2008).

During the Holocene, glaciers have fluctuated in accordance with climate changes (Denton & Karlén, 1973; Thompson et al., 2009; Solomina et al., 2015). Similar to the Heinrich events, the coldest episodes were detected in North Atlantic sea records in the form of ice-rafted debris increases known as Bond events or Bond cycles (Fig. 3; Bond et al. 1997, 2001; Mayewski et al., 2004; Wanner & Bütikofer, 2008; Isono et al., 2009; Wanner et al., 2011). Two of these climate shifts, the 8.2 and the 4.2 ka events, were registered in a variety of fossil records from all over the world (Walker et al., 2018) and provided time boundaries for the establishment of a new Holocene division into three new stages / ages (Fig. 3): the Greenlandian (11.7 – 8.2 cal ka BP), the Northgrippian (8.2 – 4.2 cal ka BP) and the Meghalayan (4.2 cal ka BP – present; Walker et al., 2018, 2019). The first millennia of the Holocene were characterised by upward temperature (Marcott et al., 2013) and humidity (Morellón et al., 2018) trends until approximately 9500 cal yr BP, coinciding with the greater Holocene values of summer insolation in the Northern Hemisphere (Fig. 3; Berger & Loutre, 1991). During that period, meltwater pulsations continued triggering episodes of cold conditions (Teller & Leverington,

2004), which were shorter than the Younger Dryas Stadial. Due to their magnitude, the 11.4, 9.3 and 8.2 ka events stand out above other pulsations (Alley et al., 1997; Alley & Ágústsdóttir, 2005; Teller & Leverington, 2004; Fleitmann et al., 2008; Rasmussen et al., 2007, 2014). In the Cantabrian region, the period around the 9.3 ka event had especially dry summers, leading to conditions similar to the Mediterranean climate (Baldini et al., 2019), while the 8.2 ka event was also a dry interval (Domínguez-Villar et al., 2009; Smith et al., 2016; Rossi et al., 2018) with the coldest winters of the Greenlandian stage (Baldini et al., 2019).

During the Northgrippian, temperature gradually decreased in southern Europe following the falling summer insolation trend in the Northern hemisphere (Berger & Loutre, 1991) and unlike the warm stable conditions exhibited at European high latitudes (Davis et al., 2003; Marcott et al., 2013). After that period, a series of cold pulsations occurred during the last millennia derived from multiple factors such as volcanic eruptions and solar irradiance fluctuations (Wanner & Buetikofer, 2008). The 4.2 ka event is well detected in North Iberian records, yielding the lowest temperature of the Holocene and summer aridity (Smith et al., 2016; Baldini et al., 2019). Thereafter temperatures fluctuated giving

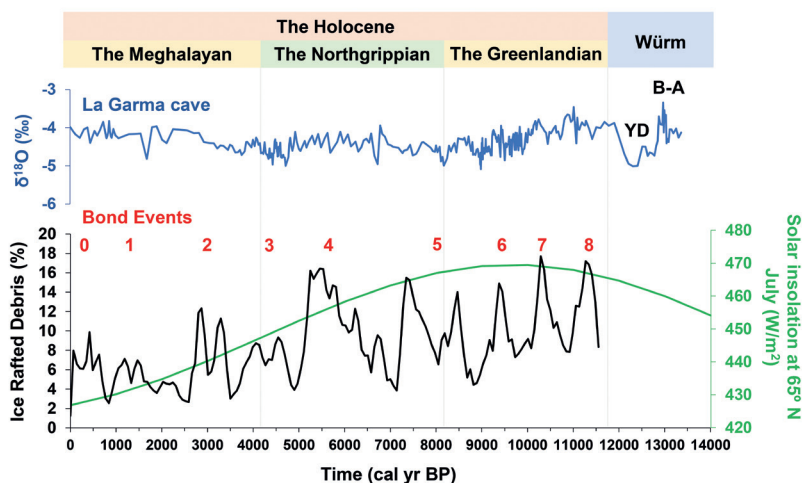


FIGURE 3. At the top: Holocene chronology based on Walker et al. (2018, 2019). In blue: hand-drilled $\delta^{18}\text{O}$ from La Garma cave speleothem (GAR-01), adapted from Baldini et al. (2019). “YD” and “B-A” refers to the Younger Dryas and to the Bolling-Allerød chronozone, respectively. Numbers in red: Bond events identified in Bond et al. (2001). Black line: Ice-Rafted Debris (IRD) index. IRD (%): Stack of MC52-V29191 + MC21-GGC22 (Adapted from Bond et al., 2001). Green line: Solar insolation of July at 65° N, adapted from Berger & Loutre (1991).

rise to several climate episodes whose chronology is well-defined on both the southern (Martín-Chivelet et al., 2011) and northern slopes of the Cantabrian range (Baldini et al., 2019). A first warm stage from 4000 to 3000 cal yr BP, including short cold pulsations in 3950, 3550 and 3250 cal yr BP, was followed by the Iron Age cold period (c. 2850 – 2500 cal yr BP) and the Roman warm period (c. 2500 – 1650 cal yr BP), both recording significant summer precipitations between 2800 and 1800 cal yr BP (Baldini et al., 2019). Following that oscillating trend, climate shifted again from the Dark Ages cold period (c. 1650 – 1350 cal yr BP, with a cold peak around 1500 cal yr BP), to the Medieval warm period (c. 1350 – 750 cal yr BP, with cold pulsations at 1250 and 850 cal yr BP) and to the last cold relapse, the Little Ice Age (LIA, 750 – 100 cal yr BP) (Martín-Chivelet et al., 2011). The 20th century started with ten of the Holocene's coldest years in Northern hemisphere mid-latitudes (Marcott et al., 2013), but temperature suddenly reversed its decreasing trend as a consequence of the greenhouse gases released into the atmosphere (Joos & Spahni, 2008), inducing the present global warming (Wanner & Buetikofer, 2008; Marcott et al., 2013).

CULTURAL DYNAMICS IN THE CANTABRIAN REGION SINCE THE PALAEOOLITHIC

Lower Palaeolithic

The oldest fossil evidence of human occupation in the Cantabrian range dates from the Lower Palaeolithic (Fig. 4). In the Sima del Elefante cave (Atapuerca cave-sites in the Atapuerca karst Mountains, southern

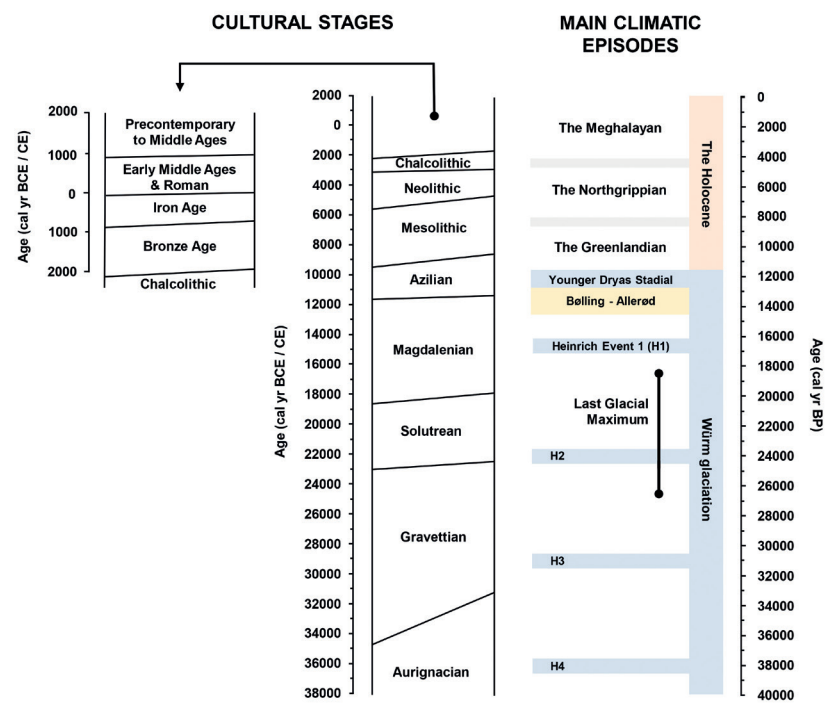


FIGURE 4. Cultural stages in Cantabria from 40,000 cal yr BP to the present and the climate characterisation of the sequence. Boundaries for the cultural stages have been drawn according to González-Sainz (1994), Rasilla-Vives & Straus (2004), González-Sainz & González-Urquijo (2004), Marín-Arroyo et al. (2018a) and Straus (2018b) for the Palaeolithic; González-Sainz & González-Urquijo (2004), Alday-Ruiz (2009), Fano et al. (2015) and Straus (2018a) for the Azilian, Mesolithic and Neolithic; Arias (1995), Ontañón (2003) and González-Rabanal et al. (2020) for the Chalcolithic; Marín-Suárez (2011) for the Bronze and Iron Ages; and Costa-García (2018) for the onset of the Roman Period. Climate: Heinrich events identified and discussed by Heinrich (1988), Bond et al. (1992) and Hemming (2004).

slope of the Cantabrian range, Burgos), some of the oldest Eurasian hominid remains are documented, dating back between 1.2 and 1.1 Myr BP (Carbonell et al., 2008). On the whole, the three main caves from the Atapuerca cave-sites, Sima del Elefante, Gran Dolina and Galería, contain hominid evidence within a range of at least 1 Myr (from 1.2 Myr to 200 ka BP; Falguères et al., 1999; Berger et al., 2008; Carbonell et al., 2008; Rodríguez et al., 2011), providing proof of the ancient presence of *Homo antecessor* and *Homo heidelbergensis* in the region. In the area equivalent to the modern Cantabrian province (hereinafter “Cantabria”), the archaeological sites from the Lower Palaeolithic date to a more recent time, most of them outdoor sites located close to the coastline (e.g., Cúlebre, Oyambre, Ubiarco, Suances, Cuchía, Hondal, Usgo, Somocuevas, Rostrío and La Verde; Fig. 5; Solórzano et al., 1999). Nonetheless, the sites with the oldest hominid remains are caves. For example, El Monte Castillo is a karst formation that hosts multiple cavities (El Castillo, La Pasiega, Las Monedas, Las Chimeneas and La Cantera caves). Up to now, El Castillo has been the site with the oldest human presence in Cantabria. A stratigraphic level, separating materials from the Acheulean and Mousterian industries, was dated at 89,000 yr BP (Bischoff et al., 1992). However, two new dates from Abrigo Rojo (close to El Mirón cave, Ramales de la Victoria, Cantabria) have recently been analysed with an estimated age of 230 – 240 ka BP, within a level that contains faunal and lithic industry remains (personal communication of Manuel R. González Morales, Universidad de Cantabria).

Middle Palaeolithic

The colder conditions of the Last Glacial Period coincided with greater occupation of caves and rocky shelters in the region by both hominids and wild animals (Solórzano et al., 1999). These cavities host multiple fossil evidence of Neanderthals and Mousterian and Châtelperronian industries (such as bones, lithic materials, phytoliths and faunal remains) as well as of their coexistence with anatomically modern humans (AMH) and Aurignacian industry. AMH arrived in Eurasia at c. 45,000 cal yr BP, their presence overlapping with Neanderthals in western Europe for a period of between 2600 and 5400 years (Higham et al., 2014). In the northern Cantabrian range, the end of the Châtelperronian is estimated at around 42.4 – 41.8 cal ka BP (Marín-Arroyo et al., 2018a). In parallel, the Aurignacian spanned from 43.3 – 40.5 to 34.6 – 33.1 cal ka BP, followed by the onset of Gravettian industry at about 36.8 – 35 cal ka BP (Marín-

Arroyo et al., 2018a). Several archaeological sites from Cantabria cover these transitional epochs (Fig. 5), although some of them are still poorly documented (Muñoz-Fernández & Llamosas, 1987; Solórzano et al., 1999). The principal stratigraphic sequences are found in caves (Table 1; Fig. 5) and in the outdoor site of El Habario (Solórzano et al., 1999).

TABLE 1. *Principal caves with stratigraphic sequences covering the Lower and the Middle Palaeolithic in Cantabria.*

Cave	References
Cobrante	Cáceres, 2009; Fernández & Santamaría, 2009.
Covalejos	Maroto et al., 2012; Cáceres, 2013; Yravedra, 2013; Yravedra et al., 2016; Marín-Arroyo et al., 2018a; Jones et al., 2019; Sánchez-Hernández et al., 2019.
El Castillo	Straus, 1975; Cabrera-Valdés, 1984, 1996; Cabrera-Valdés & Bischoff, 1989; Rink et al., 1996, 1997; Cabrera-Valdés et al., 2000, 2002; Bernaldo de Quirós et al., 2006, 2010, 2015; Sánchez-Fernández & Bernaldo de Quirós, 2008; Wood et al., 2018; Jones et al., 2019; Garralda et al., 2019; Luret et al., 2020; Martín-Perea et al., 2022.
El Cuco	Muñoz-Fernández et al., 2007; del Río et al., 2011; Gutiérrez-Zugasti et al., 2013, 2018.
El Esquilieu	Baena et al., 2005, 2012, 2019; Yravedra, 2006, 2013; Manzano-Espinosa et al., 2005; Jordá-Pardo et al., 2008; Cabanes et al. 2010; Maroto et al., 2012; Uzquiano et al. 2012; Yravedra & Castanedo, 2014; Yravedra et al., 2014; Cuartero et al., 2015.
El Mirón	Cuenca-Bescós et al., 2008, 2009; Straus & González-Morales, 2007, 2012a, 2012b, 2016; Marín-Arroyo et al., 2018b, 2020.
El Ruso	Muñoz-Fernández, 1991; Muñoz-Fernández & Serna, 1994, 1999; Yravedra et al., 2010.
El Pendo	Carballo, 1960; Echegaray, 1980; Hoyos-Gómez & Laville, 1982; Ortega-Mateos, 1982; Pike-Tay et al., 1999.
Hornos de la Peña	Ríos-Garaizar et al., 2020.
La Garma	Ontañón, 2003; Álvarez-Fernández, 2007.
Morín	Altuna, 1973; Echegaray et al., 1973; Freeman, 1978, 1983; Pike-Tay et al., 1999; Maíllo-Fernández, 2001; Maíllo-Fernández et al., 2001, 2014; Yravedra & Gómez-Castanedo, 2010; Maroto et al., 2012; Yravedra, 2013; Bradtmöller, 2015; Bradtmöller et al., 2016.

Furthermore, many of these cavities presented various artistic paintings ranging from the Middle to Upper Palaeolithic (Bicho et al., 2007), such as the decorated bone discs in the caves of El Linar and Las Aguas (de las Heras et al., 2008) and varied rock art representations in Altamira (García-Díez et al., 2013; Cuenca-Solana et al., 2016), La Garma archaeological complex (Arias et al., 2000, 2011; Arias & Ontañón, 2014; Pettitt et al., 2014) and the caves of El Castillo, Las Monedas, Las Chimeneas and La Pasiega (Fig. 5; Alcalde del Río, 1906; Alcalde del Río

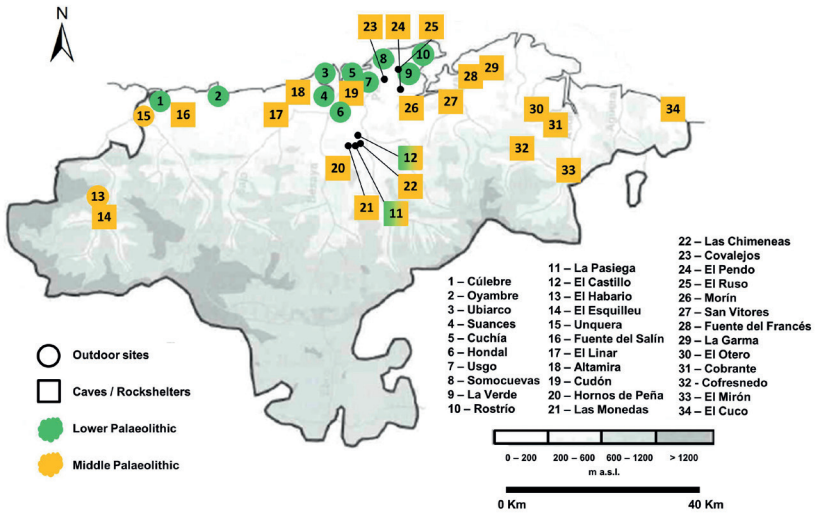


FIGURE 5. Main archaeological sites covering the Lower and the Middle Palaeolithic in Cantabria. Adapted from Solórzano et al. (1999) and completed with the sites mentioned in the text.

et al., 1912; Valladas et al., 1992; Pike et al., 2012; García-Díez et al., 2015; Pettitt et al., 2014; d’Errico et al., 2016). In particular, a U-Th dating from a carbonate crust overlying a cave painting at La Pasiega indicated a minimum age of 64.8 ka BP, which would attribute the origin of cave paintings to Neanderthals (Hoffman et al., 2018a). However, this statement has generated controversy among the scientific community and the discussion about the chronological data’s reliability and their archaeological meaning is still ongoing nowadays (Slimak et al., 2018; Hoffman et al., 2018b, 2020; White et al., 2019).

The faunal remains documented in the caves suggest that Middle-Palaeolithic Cantabrian society was mainly hunter-gatherer. Broadly, no

major differences have been detected between the diet of Neanderthals and the first *Homo sapiens* (Yravedra et al., 2016). In the case of the northern Cantabrian range, hunting bison, deer and horses seems to have been common between the Mousterian and the Gravettian, which was later joined by hunting smaller-sized ungulates such as caprids (Solórzano et al., 1999; Rasilla-Vives & Straus, 2004; Yravedra, 2013). Also, shellfish exploitation has been documented since Mousterian / Aurignacian times (Gutiérrez-Zugasti et al., 2013, 2018) and continued through the Gravettian (Álvarez-Fernández, 2007).

Upper Palaeolithic

During the Last Glacial Maximum (c. from 26.5 to 19 ka BP), the increment of the ice and permafrost cover in northern Europe (Clark et al., 2002; Vandenberghe et al., 2014; Lindgren et al., 2016) triggered massive human migrations towards southern regions (Banks et al., 2008). The Iberian Peninsula, for example, became a climate refuge (Banks et al., 2008) and Solutrean settlements increased, most of them located in coastal areas (Aura et al., 2012). More precisely, there is human evidence in regions rich in indoor sites such as caves and shelters, especially those close to watercourses to facilitate the collection of water and wood from riparian vegetation (Straus, 2015). Outdoor sites have been poorly documented (Solórzano et al., 1999), while most cavities occupied during the Gravettian continued to host human groups during Solutrean industry (Corchón-Rodríguez, 1999; Rasilla-Vives & Straus, 2004; Straus et al., 2012; Maíllo-Fernández et al., 2014; Straus, 2015; Ríos-Garaizar et al., 2020). Although the weather was cold, subsistence activities remained diversified with prey such as birds, shellfish and fish from rivers and estuaries (Rasilla-Vives & Straus, 2004; Álvarez-Fernández & Fernández-García, 2013).

During the Magdalenian, the development of lithic tools, such as harpoons in the Upper stages, was facilitated through the use of flint and derivatives which led to a further intensification of hunting (González-Sainz & González-Urquijo, 2004). In this respect, some large carnivores disappeared from the Cantabrian region due to hunting pressure (Castaños, 1992). The Red deer (*Cervus elaphus*) was the main prey during the first millennia (González-Sainz & González-Urquijo, 2004; Portero et al., 2019), especially documented in archaeological sites close to the current shoreline (<9 km) and below 200 m from the current sea level. By contrast, the Iberian ibex (*Capra pyrenaica*) was dominant at higher altitudes (Portero et al., 2019). Other food sources such as shellfish were also exploited

(Álvarez-Fernández, 2010). Additionally, important developments were made in terms of meat treatment and conservation, representing key factors in understanding the population growth (González-Sainz & González-Urquijo, 2004). Together with the warmer climate conditions, especially during the Bølling-Allerød chronozone (GI-1), communities grew up across the Cantabrian region and new cavities were occupied (Fig. 6; González-Sainz, 1989; González-Sainz & González-Urquijo, 2004; Chauvin, 2007; Chauvin et al., 2018; Fano et al., 2020). In particular, human groups expanded towards mid-mountain areas (0 – 700 m a. s. l.) along the Upper Magdalenian and the first stages of the Azilian, achieving increasing control of landscape with important migration rates among sites (Fernández-Tresguerres, 2004; González-Sainz & González-Urquijo, 2004). Economic activities between the Cantabrian region, the Pyrenees and south-western France became frequent, as evidenced by the exchange of art pieces, tools and several valuable objects (Sauvet et al., 2008; Straus & Langlais, 2020; Lefebvre et al., 2021). Cave art was also particularly relevant during the Magdalenian (González-Sainz, 2007, 2012), especially because the paintings with the most recent radiocarbon dates belong to that period (three radiocarbon dates of paintings between 14,000 and 13,600 cal yr BP in Las Monedas cave: González-Sainz, 2005).

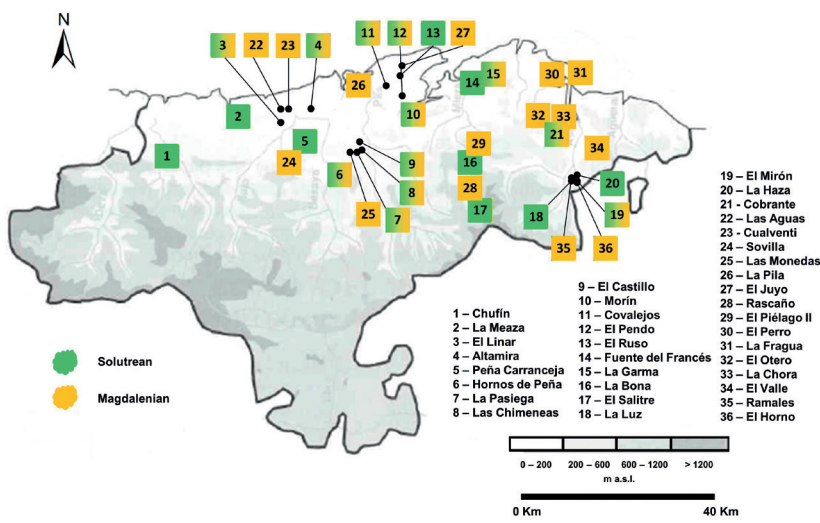


FIGURE 6. Main archaeological sites covering the Solutrean and Magdalenian cultural stages in Cantabria. Adapted from Solórzano et al. (1999) and completed with sites described in Rasilla-Vives & Straus (2004) for the Solutrean and González-Sainz & González-Urquijo (2004) for the Magdalenian.

Azilian and Mesolithic

The Azilian refers to the Epipalaeolithic culture developed in several eastern sites in the northern Cantabrian range among other European regions. It encompassed the second half of the Bølling-Allerød chronozone ($\approx$ since the Greenland Interstadial 1c), the Younger Dryas Stadial (GS-1) and the first stages of The Holocene. Technologically, the decrease in tool size and variety represented an important shift with respect to the Magdalenian (Fernández-Tresguerres, 2006), and the reduction in artefacts made of bones is especially significant (Fernández-Tresguerres, 2004). However, small-size stone tools, also known as microliths, started to spread during that time. The diagnostic tool of the period was a flat harpoon instead of the circular harpoon characteristic of the Upper Magdalenian. Thus, human populations still based their economies on hunter-gatherer activities (Fernández-Tresguerres, 2004). Human presence was detected at high altitudes of the northern Cantabrian range, for example in the Portugain rock shelter (920 / 940 m a. s. l., Navarra) and the Urratxa III cave (1015 m a. s. l., Vizcaya; Barandiarán et al., 2006). However, there are no documented sites within this altitude range in Cantabria. Moreover, similar to the decrease in tool variety, artistic expression was also drastically reduced with only a limited number of art pieces in comparison to earlier ages (Fernández-Tresguerres, 2006). Altogether, however, all these changes seem to have come about progressively, while the Younger Dryas Stadial does not appear to have marked a clear boundary in cultural progression (Straus, 2011, 2018a).

The microlith industry continued to grow on the eastern side of the northern Cantabrian range during the Mesolithic (Alday-Ruiz & Cava, 2009; Soto et al., 2015), while a new cultural phase, the Asturian, was settled on the most westerly side (there are only a few Asturian sites on the most westerly side of Cantabria). In Cantabria, 256 non-Asturian sites have been documented, the largest concentration of which was found close to the coastline and along the littoral plain, although some interior valleys were also frequented. The maximum altitude was achieved by three sites at between 650 and 750 m a. s. l. (Pérez-Bartolomé, 2019).

In terms of climate, temperature followed an upward trend during these millennia (Marcott et al., 2013). Accordingly, the sea level continued rising as it had done since the Last Glacial Maximum (Spratt & Lisiecki, 2016), which would have progressively affected some ancient migratory routes between Cantabria, the Basque country and France (Straus, 2018b). Also, the 9.3 and 8.2 ka climate events occurred during the Mesolithic,

both shorter than the Younger Dryas Stadial. The 9.3 ka event brought very dry summers in Cantabria while during the 8.2 ka event winters were especially harsh (Baldini et al., 2019). Considering the archaeological records, however, these events do not seem to have led to significant social changes in the region (Straus, 2018a), contrary to what happened in dryer areas in the Mediterranean basin (Berger & Guilaine, 2009). In the Iberian Peninsula, for example, human populations migrated from the Lower Aragón to more humid northern areas during the centuries after the 8.2 ka event, leaving an archaeological silence with no known human evidence until the Neolithic period (González-Sampériz et al., 2009).

Neolithic

One of the most striking aspects of the Neolithic period was the transition from hunter-gatherer to farmer societies (Arias, 1991) which took place along the Cantabrian range at between 7700 – 6800 cal yr BP (Fano et al., 2015). There seems to be evidence that these activities were introduced by foreign human groups, although it remains unclear whether the migrations came from the Pyrenees (Zapata, 2000; Zapata & Peña-Chocarro, 2005) or the Ebro Valley (Arias, 2007). In this respect, radiocarbon dates place the oldest cereal remains from the northern Cantabrian range at between c. 7000 and 6000 cal yr BP (Table 2), some centuries after the first documented evidence in Iberian Mediterranean sites (Alday-Ruiz, 2009; Pérez-Obiol et al., 2011). In particular, up to four different cereals have been identified in the region, namely einkorn, emmer, barley and wheat (Zapata et al., 2004), while their concentration was lower than in other Iberian sites (Straus, 2018a). Likewise, the domestication of animals for human consumption went hand in hand with the onset of agriculture. The oldest domesticated fauna remains date approximately from 6900 to 6500 cal yr BP (Cubas & Fano, 2011). Among other species, cows, pigs and ovicaprid fauna have been identified as part of the diet composition of Cantabrian Neolithic and Chalcolithic human groups (Arias, 2005; Aranburu-Mendizabal et al., 2018; Sarasketa-Gartzia et al., 2018; Jones et al., 2019; González-Rabanal et al., 2020). However, ancient activities such as animal hunting and gathering of marine resources and wild plants persisted significantly after the introduction of agricultural practices in many settlements (Cubas et al., 2016). Thus, it is believed that both traditional and farmer human groups lived together for at least 500 years in the Cantabrian region before the establishment of more developed agricultural societies (Fano et al., 2015). Additionally and also related

to the latter developments, ceramic utensils spread during the Neolithic (Alday-Ruiz, 2009). Their origin is probably linked to farming activities as they may have been used as storage for surplus food or as containers to facilitate trading among human groups (Orton et al., 1997).

As a whole, the Neolithic archaeological record may seem limited in Cantabria with only a few documented sites, such as the caves of El Mirón (Straus & González-Morales, 2012a, b), Los Gitanos (Ontañón et al., 2013) and La Calvera Hut (Diez-Castillo, 1996). However, a funerary ritual introduced in the form of megaliths reveals intense anthropic occupation. In particular, approximately 1250 megaliths have been documented in the northern Cantabrian range, distributed from sea level to more than 1800 m a. s. l. (Arias et al., 2005). The propagation of these funerary monuments occurred rapidly after their introduction, which is dated at around 6550 – 6200 cal yr BP (Fano et al., 2015).

To conclude, all these socioeconomic breakthroughs triggered changes in food resource exploitation and yielded more cohesive human communities, resulting in the appearance of prominent settlements around Europe (Sherratt, 1995).

TABLE 2. *Documented cereal evidence from the modern provinces of the Northern Cantabrian range (Asturias, Cantabria and the Basque Country).*

Region	Site	Type of evidence	Oldest documented age (cal yr BP)	Reference
Basque Country	Lumentxa	Seed	6178 – 5747	Zapata (2002)
	Kobaederra	Seed	6658 – 6219	Zapata (2002)
		Seed	6308 – 5938	Zapata (2002)
	Pico Ramos	Seed	6280 – 6004	Zapata et al. (2004)
	Herriko	Pollen	7156 – 6661	Iriarte-Chiapusso et al. (2005)
	Barra	Pollen	7153 – 6539	Iriarte-Chiapusso et al. (2005)
	Los Husos	Pollen - Seed	7172 – 6849	Iriarte-Chiapusso (2009)
Cantabria	El Mirón	Seed	6484 – 6016	Peña-Chocarro et al. (2005)
		Seed	6821 – 6357	Peña-Chocarro et al. (2005)
		Seed	6406 – 6287	Peña-Chocarro et al. (2005)
	La Molina	Pollen	6735 – 6495	Pérez-Obiol et al. (2016), Sánchez-Morales et al. (2022)
Asturias	Monte Areo	Pollen	6736 – 6495	López-Merino et al. (2010)

Chalcolithic, Bronze and Iron Ages

As noted by Blanco-González et al. (2018), later Prehistory has been less studied than the Palaeolithic period in the Cantabrian range, which also started after it did in the Mediterranean basin. However, several significant new factors are worth emphasising.

During the Chalcolithic, human presence spread over the entire Cantabrian region, with a geographical level of organisation comprising a similar extension to the modern Spanish *comarcas* (Ontañón, 2015). Economic activities continued to be based on agricultural and farming practices (Ontañón, 2015), thus solidifying the model introduced during the Neolithic with increasing anthropic pressure on the landscape. Evidence of domestic animals increased during this period (Ontañón, 2003; Altuna & Mariezkurrena, 2009), in particular cattle and pigs, as documented in sites such as El Mirón cave (Altuna & Mariezkurrena, 2012) and La Castañera rocky shelter (Vega-Maeso et al., 2016). Although hunting of wild animals decreased significantly (Ontañón, 2000), traditional activities did not disappear. Thus, Cantabrian Chalcolithic industry produced a great quantity of lithic projectile points. Their function varied since they could well be used to hunt, forming part of the arrows of the remaining hunter groups, or also employed as offerings in funerary rituals (Ontañón, 2002). These artefacts have been found around megaliths, which continued to be important funerary elements, but also in caves (Ontañón, 2002). This is due to the fact that caves and shelters started to be used as burial sites, which in turn went hand in hand with the progressive decrease in human occupation (Ontañón & Armendáriz, 2005). Along with many other sites, the Los Avellanos I and II caves (Cantabria) are examples of these funerary rituals (González-Rabanal et al., 2020).

As a whole, all changes mentioned above should be framed within the scenario of climate stability. The most relevant oscillation took place during the 4.2 ka event, which did not significantly affect anthropic activity (Balsera et al., 2015; Blanco-González et al., 2018).

By the end of the Chalcolithic, industry shifted as metal pieces started to be developed. Although copper (Cu) deposits were irregular and scarce in Cantabria (Mantecón, 2000), metal exploitation was particularly important on the westernmost side of the Cantabrian range. For example, the Aramo range (Asturias) provided great quantities of copper (Blas-Cortina, 2014). Thus, the first metal utensils consisted of this element, which was alloyed with arsenic (As) by the first half of the second millennium BCE (Early Bronze Age). After that, copper was alloyed with

other materials in the Bronze Age, first with tin (Sn) to obtain Bronze, and thereafter with tin and lead (Pb), albeit ternary alloys seem to have been poorly developed in Cantabria (Arias et al., 2005). In parallel, the first fortified structures of the Iberian Peninsula, which are in some cases referred to as Castros, were built during the Chalcolithic. However, it was not until the Late Bronze Age when the first fortifications were erected in the Cantabrian region (e.g., Castro de la Garma dates from 1400 BCE: Obregón-Goyarrola, 2010).

Contrary to the case of copper, Cantabria hosted several iron mines, especially in its central and eastern sectors (Mantecón, 2000). The development of iron metallurgy led to a diversification of weapons and warlike equipment (Torres-Martínez, 2010). Also, fortified structures became more numerous during the Iron Age and the Roman period when wars and battles started to take place between peoples. In this respect, most fortifications were placed at high altitudes to have good visibility over and control of the surrounding territories. Thus, settlement distribution shifted until the Middle Ages, when they were again placed along interior valleys. The modern comarca of Campoo, with an average altitude above 1000 m a. s. l., is the Cantabrian area with more documented Castros (Obregón-Goyarrola, 2010).

From the Roman period to the present

In Cantabria, the end of the Iron Age is delimited by the invasion of the Roman Empire, which took place after a series of battles between 29 and 19 BCE known as the Cantabrian Wars (Costa-García, 2018). The Roman troops entered the Cantabrian region from the southern slope of the Cantabrian range divided into two main battlefronts with the aim of taking control of Noega (Gijón, Asturias) and Portus Victoriae (Santander, Cantabria). In both cases, mountain areas became the battlefield of the wars where a great number of roads were also built to facilitate the advance of troops (Peralta-Labrador et al., 2019).

The process of Romanisation was a culture shock in Cantabria as it brought a new hierarchical social system and gender inequality which has not been detected in previous ages (Marín-Suárez & González-Álvarez, 2011). Otherwise, traditional mining works continued to be important during that time, with a total of 10 Roman mines located along the central sector of the littoral plain (Fig. 7). The extractions were principally of iron (hematite and goethite ores), lead (galena and cerussite, also containing a

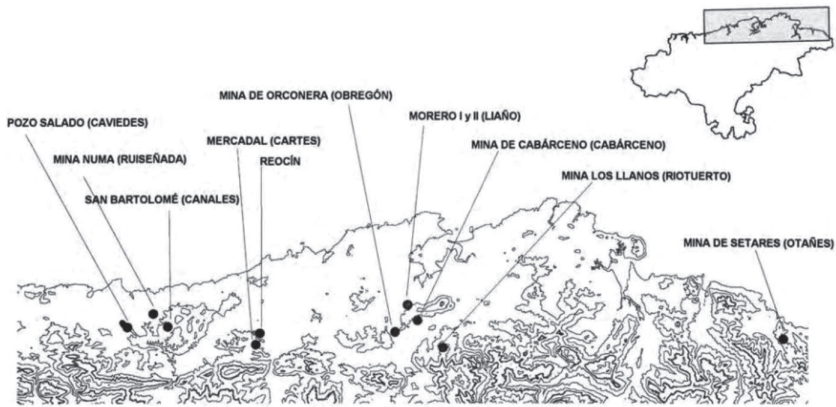


FIGURE 7. *Location of Cantabrian mines exploited during the Roman period. Retrieved from Mantecón (2000).*

variable proportion of silver which was difficult to separate), calamine and probably some salt mines which were exploited as well (Mantecón, 2000). Gold mines, by contrast, were more abundant in Asturias and Galicia (Lewis & Jones, 1970; Spiering et al., 2000).

The Roman Empire began to lose its power in the early 400s CE. Two Germanic groups, the Sueves and the Vandals, together with the Alans, a people from further east (probably Iranian), crossed the Pyrenees in 409 CE and took control of some Iberian points. The Visigoths, a Germanic group allied to the Romans, regained territory after fierce battles during the 5th and 6th centuries and established the Visigoth Kingdom in the Iberian Peninsula (Collins, 2004). Afterwards, the Middle Ages brought a Christian monarchy to Cantabria for the first time, the Kingdom of Asturias, which encompassed the north-western part of the Iberian Peninsula. The most prominent buildings of the Middle Ages were castles that were deployed around various areas of Cantabrian territory depending on the period. The oldest were constructed on mid-mountain areas, while between the 8th and 11th centuries CE the fortifications returned to high altitudes, similar to the Iron Age distribution. Around the 14th and 15th centuries, however, castles were mostly sited in low valleys, close to villages or in them (Martínez & Mantecón, 2012).

In terms of economic activities, maritime industry gained importance in Cantabria during the Middle Ages, which led to a greater need for iron and wood. The importance of this industry was such that the Crown of Castile gave preferential treatment to operations in the Cantabrian Mountains in 1335 CE (Corbera, 1999). Here major breakthroughs were

made by using charcoal as fuel instead of freshly-cut wood during the 12th-13th centuries CE, resulting in a calorific power increase (García-Alonso, 1999). In Europe, the first hydraulic finery forges date from the 12th century CE. Although it is estimated that they were introduced into the Iberian Peninsula during the 13th century CE it was not until the 14th century CE that they are documented in Cantabria closely linked to the shipbuilding industry, and increased significantly during the 15th century CE (Corbera, 1999; Cuerno, 2001). The time of maximum charcoal demand, however, is documented as from 1622 CE, when a cannon industry was constructed in the towns of Liérganes and La Cavada (García-Alonso, 1999). Charcoal kilns were widely scattered across the territory by that time, resulting in significant levels of deforestation (García-Alonso, 1999).

During the 17th century CE, a socioeconomic crisis hit Spain and also Europe on a larger scale. The iron industry withstood the crisis owing to its importance until 1850 – 1870 CE, when iron production decreased in Cantabria due to the depletion of forest resources and the boom in foreign sources of supply (Corbera, 1999). In accordance, the use of charcoal kilns progressively decreased during the middle of the 19th century CE (García-Alonso, 1999).

THE STUDY OF THE CANTABRIAN LANDSCAPE'S EVOLUTION BASED ON PALEOLIMNOLOGY

Due to the great variety of data, multiproxy paleolimnological studies have become more numerous in recent times, emerging as useful tools to better characterise past environmental dynamics (Birks & Birks, 2006; Michelutti & Smol, 2013). The number of studies in the north-western Iberian Peninsula has steadily increased over the last century, both in the Atlantic and Mediterranean biogeographical regions (Fig. 10; Table 3). This has been possible owing to the large number of glacial lakes (Jiménez & Farias, 2005) and peat deposits (Martínez-Cortizas et al., 2000, 2009; Ramil-Rego et al., 2018) in the area. The sequences with the longest chronology in the Iberian north-west, however, are the ones obtained from two marine records (MD03-2697 and MD99-2331) which were sampled in the Atlantic Ocean and have a bottom age of 340 ka BP (Gouzy et al., 2004; Sánchez-Góñi et al., 2005, 2008; Naughton et al., 2007; Desprat et al., 2009).

In the Cantabrian range, the reconstruction of plant dynamics has been enabled by the large number of palynological records studied (Fig. 10; Table 3). From west to east, the sequences of Laguna de Lucenza (Aira-Rodríguez, 1986; Santos-Fidalgo et al., 1997; Muñoz-Sobrino et al., 2001), Villaseca de Laciana (Jalut et al., 2010), La Mata (Jalut et al., 2010), Lago de Ajo (Allen et al., 1996), Puerto de Tarna (Ruíz-Zapata et al., 2000), Puertos de Ríofrio (Menéndez-Amor & Florschütz, 1963) revealed an open area with a large presence of xerophytes during some intervals of the Last Glacial Period above 1000 m a. s. l. The same landscape was inferred on the northern slope of the Cantabrian range from the record of La Molina peat bog (Cantabria, 484 m a. s. l., Sánchez-Morales et al., 2022). A few more records from Burgos (north-eastern Castile and León), such as La Piedra (Ramil-Rego et al., 1998) and San Mamés de Abar (Iriarte-Chiapusso et al., 2001), documented a higher presence of pines and birches, while the dominance of pines becomes more evident in the northern Iberian System as indicated by the papers on Hoyos de Iregua (Gil-García et al., 2002), Laguna del Hornillo (Gómez-Lobo et al., 1996), Laguna Negra (Von Engerlbrecthen, 1998), Lago Las Pardillas (Sánchez-Góñi & Hannon, 1999), Quintanar de la Sierra (Peñalba, 1994; Peñalba et al., 1997) and Laguna Grande (Ruiz-Zapata et al., 2002a).

On the northern slope of the Cantabrian range, plant dynamics studies have revealed a generalised dominance of mixed deciduous formations over the Holocene epoch composed of deciduous *Quercus*, *Corylus* and *Betula* at some stages. In Asturias, the papers on El Alto de la Espina (López-Merino,

2009), Las Dueñas (López-Merino et al., 2006), El Monte Areo peat bogs (López-Merino, 2009; López-Merino et al., 2010) and Lake Enol (López-Merino, 2009) are the most relevant. In Cantabria, located in the eastern sector of the mountain system, the main sequences are those obtained from the peat bogs of Culazón (González-Pellejero et al., 2014), Pico Sertal (Mariscal, 1986; Carracedo et al., 2018), El Cueto de la Avellanosa (Mariscal, 1983; Núñez, 2018), Pico Ano (Salas, 1993), Alsa (Mariscal, 1993), Puerto del Escudo (Muñoz-Sobrino, 2001), Estacas de Trueba (Mariscal, 1987, 1989), Sotombo (Pérez-Díaz et al., 2016a), Los Tornos (Peñalba, 1994; Muñoz-Sobrino et al., 2005), La Nava (Menéndez-Amor, 1968) and La Molina (Pérez-Obiol et al., 2016; Sánchez-Morales et al., 2022). Some syntheses containing most of these paleobotanical records and also archaeological sites from the northern Iberian Peninsula were provided by Iriarte (2009) and Pérez-Obiol et al. (2011). Other paleobotanical syntheses have also been performed for the Western Pyrenees region (Pérez-Díaz et al., 2015) and for the Iberian Peninsula and Balearic Islands (Carrion, 2012).

Among all the mentioned sites, however, only a few studies considered the role of fire in the reconstruction of the Cantabrian landscape. This may lead to partial interpretations since fire is an important factor in the complex scenario between vegetation, climate and anthropic pressure (Harrison et al., 2010; Whitlock et al., 2010; Krawchuk & Moritz, 2011). The fire regime can be characterised by analysis of sedimentary charcoals. Charcoals are plant remains resulting from incomplete combustion (Whitlock & Larsen, 2001) that can be transported by wind over long distances (Radtke et al., 1991; Andreae, 1991). Some studies estimate that the smallest particles (<100 µm) can travel more than 100 km through the atmosphere (Conedera et al., 2009), while charcoal particles larger than 125 µm are mainly deposited within a 7-km radius (Whitlock & Millspaugh, 1996). Therefore, the study of sedimentary micro and macrocharcoals can provide insight into the fire regime on both a regional and local scale respectively. In Europe, analysis of multiple charcoal sequences allowed the determination of fire episodes that occurred simultaneously in different regions during some glacial and postglacial intervals and therefore might have been strongly influenced by climate conditions (Power et al., 2008; Daniau et al., 2010, 2012; Marlon et al., 2016). In the Cantabrian range, the oldest charcoal record is the obtained from La Molina peat bog, that covers the last c. 17,500 cal years (Sánchez-Morales et al., 2022). The lack of woody species during the Palaeolithic section of the record went hand in hand with the absence of significant fire events at a local scale, in accordance with other Iberian paleosites (Portugal: Connor et al., 2012; Central Pyrenees: Gil-Romera et al., 2014; Andalucía: Carrión, 2002). The first significant fire episode occurred during the early Holocene

coinciding with the 9.3 and the 8.2 ka events (Sánchez-Morales et al., 2022). “It has been shown that climate changes of this magnitude can cause major fires in a brief period of time (Daniau et al., 2019) since both temperature and humidity are acknowledged to be important factors in controlling biomass burning on a regional scale (Daniau et al., 2012). Power et al. (2008) also point out that increased seasonality during the early Holocene could have regulated the fire regime in the Northern Hemisphere” (Sánchez-Morales et al., 2022: 107373). Thus, the dry conditions of the 9.3 and the 8.2 ka events in the Cantabrian range (Domínguez-Villar et al., 2009; Smith et al., 2016; Rossi et al., 2018; Baldini et al., 2019) together with the fuel accumulation that took place during the early Holocene (Sánchez-Morales et al., 2022) were most likely responsible of the two fire episodes.

For the last 7000 years, the sequences obtained from El Sertal, El Cueto de la Avellanosa (Carracedo et al., 2018) and La Molina (Pérez-Obiol et al., 2016; Sánchez-Morales et al., 2022), all located in the eastern sector of the Cantabrian range, indicated intense use of fire from the Neolithic onwards coinciding with the onset of agricultural practices and the need to create forest openings, thus revealing an anthropic origin (Figs. 8 and 9). In La Molina, which is the study site located at the lowest altitude (484 m a. s. l.), important fire events were detected from 5800 to 3500 cal yr BP (Late Neolithic, Chalcolithic and Bronze Age) in the form of intense and frequent charcoal peaks and increases in charcoal content. The beginning of these fire events was accompanied by the appearance of *Calluna* and other Ericaceae (Pérez-Obiol et al., 2016) whose germination is stimulated by high temperatures (Reyes & Casal, 2000). Fire intensity was lower at a higher altitude range, especially in El Sertal (940 m a. s. l.) where sedimentary macrocharcoals were not continuously recorded until the Middle Ages, increasing sharply during the last 400 years. By contrast, in El Cueto de la Avellanosa (1320 m a. s. l.) the fire signals were constant from the beginning of the sequence, registering a rise at around 3600 cal yr BP.

Paleolimnology has also been applied in the Cantabrian region to determine the beginning and the main phases of former mining activities. This has been possible since the atmospheric pollution resulting from mining or metallurgy, as well as from other anthropic sources such as the combustion of coal or petrol and waste incineration, can result in trace metal (e.g., Pb, Hg, Fe, Zn) release and subsequent dumping in sedimentary deposits (Boyle, 2001). However, some caution is needed in the interpretation of trace metals since they may also come from other sources such as erosional processes and can be remobilised within the sedimentary layers due to redox conditions (Boyle, 2001). In this respect, it is interesting to look at

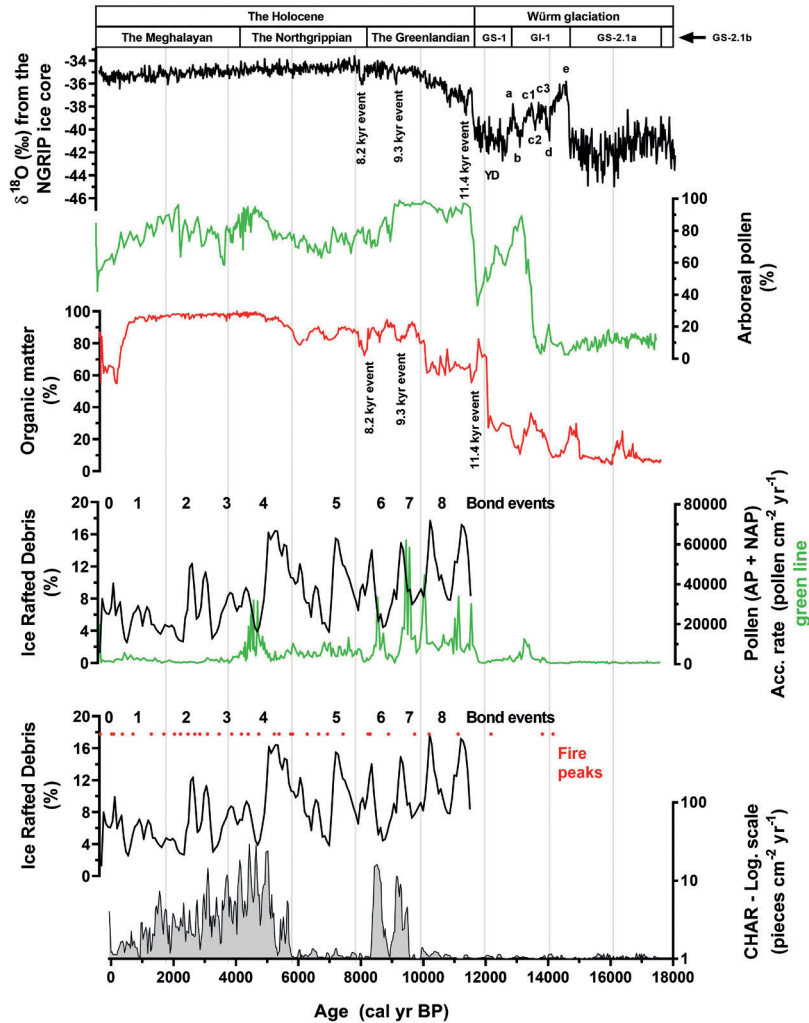


FIGURE 8. Comparison of CHAR, Ice Rafted Debris (IRD) index: Stack of MC52-V29191 + MC21-GGC22 (Bond et al., 2001), pollen accumulation rate (arboreal: AP + non arboreal: NAP) and organic matter of La Molina peat bog (Sánchez-Morales et al., 2022) and $\delta^{18}\text{O}$ from the NGRIP ice core, adapted from Rasmussen et al. (2014). At the top of the figure, the chronology for the Würm glaciation based on Greenland Stadials (GS) and Greenland Interstadials (GI) is shown (Rasmussen et al., 2014), and also the stages for the Holocene based on Walker et al. (2018, 2019). The Younger Dryas Stadial (YD), Heinrich event 1 (H1; Heinrich, 1988; Bond et al., 1992; Hemming, 2004) and the substages of the Greenland Interstadial 1 (a-e) are indicated. Retrieved from Sánchez-Morales et al. (2022).

the isotopic composition of Pb. This is a toxic element found in multiple anthropic emissions whose composition of stable isotopes varies depending on the primordial Pb, U and Th concentrations of the source and the lengths of the decay processes (Komárek et al., 2008). ^{204}Pb is a primordial lead isotope whose concentration on Earth does not vary over time, whereas ^{206}Pb , ^{207}Pb and ^{208}Pb isotopes are products of the radioactive decay of ^{238}U , ^{235}U and ^{232}Th respectively (Long, 1999). Therefore, each material containing Pb has a different isotopic signature and thus its exploration in sediments can provide a fine-tuned view of the sources of metal inputs.

In Asturias (north-western Cantabrian range), the study of a sedimentary sequence from La Molina mire detected evidence of mining activities from the beginning of the Chalcolithic (c. 4980 cal yr BP) in the form of important concentrations of trace metals and changes in lead isotopic composition (Martínez-Cortizas et al., 2013, 2016). In Cantabria former exploitations of iron and zinc ore deposits have also been documented, but to date there are no paleolimnological studies that have explored the sources and scale of the mining phases.

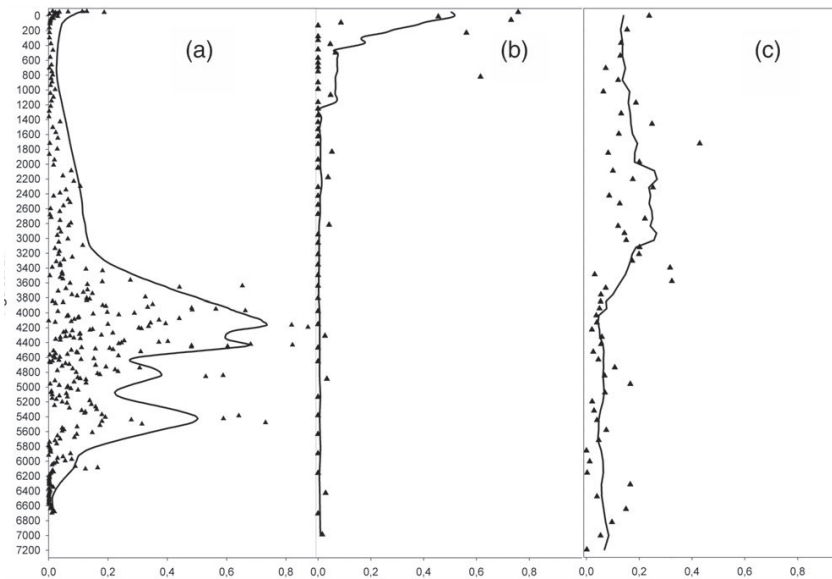


FIGURE 9. Comparison of the charcoal accumulation rate (CHAR) with rescaled and smoothed values from La Molina (a), El Sertal (b) and El Cueto de la Avellanosa (c). Over the past 7000 years, the intensity and chronology of fire events were unequal at the three sites studied, which can be interpreted as the result of asynchronous human activity in the high versus low mountain areas and also of the type of fuel involved in each zone. Retrieved from Carracedo et al. (2018).

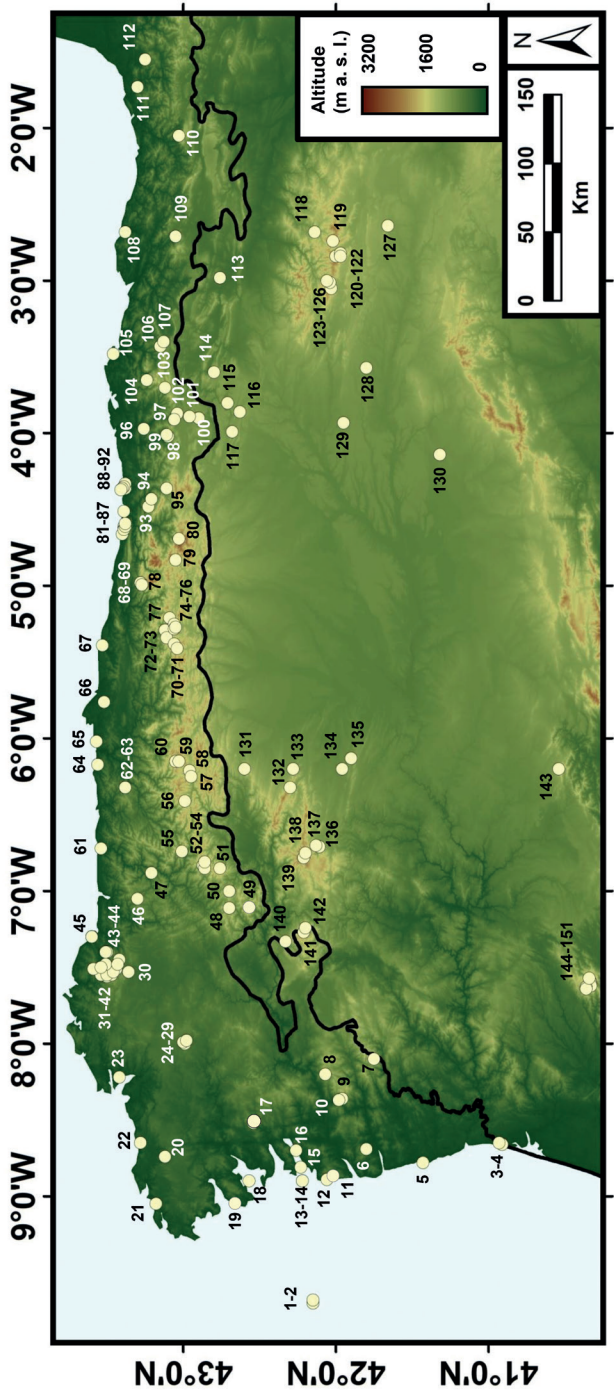


FIGURE 10. Map of the sedimentary sites studied in the north-western Iberian Peninsula. The black line indicates the border between the Eurosiberian and Mediterranean biogeographical regions (northern and southern sides, respectively).

Table 3. Sedimentary sites from the north-western Iberian Peninsula.

Code	Site	Region	Lat.	Long.	Altitude (m a. s. l.)	Estimated chronology of paleontological data (ka BP)	Palynology	Plant macroremains	Sedimentary microcharcoals	Organic matter	Metal content / Pb isotopes	Other paleodata	References
1	Marine record: MD03-2697	Atlantic ocean	42.15	-9.70	0	c. 340 to 270 // c. 12.5 to present c. 135 to 13.2	X					X	Naughton et al. (2007), Desprat et al. (2009), redrawn in Carrión et al. (2012)
2	Marine record: MD99-2331	Atlantic ocean	42.15	-9.68	0		X					X	Gouzy et al. (2004), Sánchez-Gómez et al. (2005, 2008), Naughton et al. (2007), redrawn in Carrión et al. (2012)
3	Maceda	Portugal	40.91	-8.66	0	c. 25 to 15	X						Gómez-Orellana et al. (2001)
4	Cortegada	Portugal	40.93	-8.65	0	c. 6.8 to 5.5	X						Gómez-Orellana et al. (2001)
5	Aguaçadoura	Portugal	41.43	-8.78	0	c. 4.4 to 0.3	X						Gómez-Orellana et al. (2001)
6	Serra de Água	Portugal	41.80	-8.69	780	c. >7.1 to present	X						Gómez-Orellana et al. (2010)
7	Lagoa do Marinho	Portugal	41.75	-8.1	1150	c. >11.2 to present	X						Ramírez-Rego et al. (1998)
8	Ribo Moador: PUT V	Portugal	42.07	-8.20	970	¹⁴ C data at 3.4	X						Ruiz-Zapata et al. (1995)
9	Lortelro-Cha do Couço: PUT V	Portugal	41.96	-8.36	500	¹⁴ C data at 3	X						Ruiz-Zapata et al. (1995)
10	Portela de Alvião: PUT III	Portugal	41.98	-8.37	480	Ages not provided	X						Gómez-Orellana et al. (1995)
11	Santa Maria de Oia	Pontevedra	42	-8.87	3	¹⁴ C dates at 34.2 and 29.3	X						Saa-Otero & Vías-Fleitos (1988), Gómez- Orellana et al. (1998), Ramil-Rego et al. (1998)
12	Mougas lagoon	Pontevedra	42.06	-8.89	0	c. 18.8 to present	X						Costas et al. (2009)
13	Islas Cíes: CS5	Pontevedra	42.22	-8.90	0	c. >5.2 to <3.7	X						Orellana et al. (2009)
14	Islas Cíes: CS9	Pontevedra	42.22	-8.90	0	¹⁴ C data at 3.4	X						Costas et al. (2009)
15	Ria de Vigo: MVR-3	Pontevedra	42.23	-8.81	30	MIS 3 // c. 10.4 to present	X						Muñoz-Sobrino et al. (2018), García-Moreiras et al. (2019a)
16	Ria de Vigo: B5	Pontevedra	42.26	-8.70	18	MIS 3 // c. 8.5 to present	X						Muñoz-Sobrino et al. (2018), García-Moreiras et al. (2019a)
17	Campo Lameiro area - Monte Paradela	Pontevedra	42.53	-8.51	310-260	c. 9.8 to present	X	X					Kaal et al. (2011, 2013), López-Merino et al. (2012)
18	Ria de Arousa: A14-VC15	Pontevedra - A Coruña	42.57	-8.9	0	c. 12.5 to 8.7	X					X	García-Moreiras et al. (2019b)
19	Caamaño A and B	A Coruña	42.6	-9.02	3	c. 36 to <28.7 // ¹⁴ C data at 20.2	X						Gómez-Orellana et al. (2001, 2013)
20	Bría Rubia	A Coruña	43.12	-8.74	390	c. >2.5 to present	X						Törnqvist et al. (1989)
21	Troada	A Coruña	43.18	-9.05	4	No paleontological data		X					Beo et al. (2007)
22	Baldayo	A Coruña	43.28	-8.95	0	No paleontological data							García-Amorén et al. (2008)
23	Seiselle	A Coruña	43.42	-8.92	0	c. 4.3 to present	X						Santos-Piñero et al. (1993)
24	Cruz de Bocelo	A Coruña	42.96	-9.00	745	c. 2.1 to present	X						Castro et al. (1993)
25	Insua	A Coruña	42.96	-9.00	750	c. 3.0 to present	X						Castro et al. (1993)
26	Munio	A Coruña	42.99	-7.99	705	c. 2.6 to present	X						Aira-Rodríguez et al. (1994), Castro et al. (1993)
27	Aneninos	A Coruña	42.99	-7.99	700	c. 6.5 to present	X						Castro et al. (1993)
28	A Lagoa	A Coruña	42.98	-7.98	710	c. 3.1 to present	X						Castro et al. (1993)
29	A Pena	A Coruña	42.98	-7.98	660	c. 0.16 to present	X						Ramil-Rego & Aira-Rodríguez (1993c)
30	Sever	Lugo	43.36	-7.53	620	c. 5.1 to present	X						

Table 3. Continued.

Code	Site	Region	Lat.	Long.	Altitude (m a. s. l.)	Estimated chronology of palynological data (ka BP)	Palynology	Plant macroremains	Sedimentary microcharcoals	Sedimentary microcharcoals	Organic matter	Metal content / Pb isotopes	Other paleodata	References
31	Pena Veira	Lugo	43.42	-7.45	585	c. 5.5 to present	X							Ramíl-Rego & Aíra-Rodríguez (1993b)
32	Tremolal de Schwejk	Lugo	43.43	-7.48	600-650	Ages not provided	X							Ramíl-Rego & Aíra-Rodríguez (1994)
33	Pena Vella	Lugo	43.44	-7.51	700	c. >9.6 to present	X							Ramíl-Rego et al. (1994, 1998)
34	Tremolal do Pedrido	Lugo	43.45	-7.53	695	No palynological data						X	X	Old et al. (2010), Pérez-Rodríguez & Martínez-Cortizas (2014)
35	Charco do Chan da Cruz	Lugo	43.46	-7.53	800	c. 5.9 to present	X							Ramíl-Rego & Aíra-Rodríguez (1993a)
36	Tremolal do Río das Furnas II	Lugo	43.47	-7.55	800	¹⁴ C data at 7.4	X							Ramíl-Rego et al. (1998), redrawn in Muñoz-Sobrinho (2001)
37	Turbera Pena da Cadela	Lugo	43.50	-7.55	970	c. >4.6 to present	X	X				X	X	Martínez-Cortizas et al. (2002, 2007), Mighall et al. (2006), Old et al. (2010), Castro et al. (2015), Souto et al. (2016)
38	Turbera de Chan do Lamoso	Lugo	43.50	-7.56	1150	c. >8.8 to present	X					X		Ramíl-Rego et al. (1998), Muñoz-Sobrinho et al. (2005), Old et al. (2010)
39	Turbera de Penido Vello	Lugo	43.54	-7.55	700	c. >3.4 to present	X					X	X	Martínez-Cortizas et al. (1997, 1999, 2002), Muñoz-Sobrinho et al. (2005), Kylander et al. (2003), Old et al. (2010), Rauch et al. (2010), Schellekens et al. (2011, 2015), Pontevedra-Pombar et al. (2012)
40	Borrallieras de Cal Grande	Lugo	43.59	-7.51	600	c. >4.6 to present	X	X						Mighall et al. (2006), Souto et al. (2016), Souto (2016)
41	Chao de Veiga Mol	Lugo	43.54	-7.50	700	No palynological data		X					X	Souto et al. (2016), Pontevedra-Pombar et al. (2019), Castro et al. (2020)
42	Gañidoira	Lugo	43.51	-7.48	720	¹⁴ C data at 6.9	X							Ramíl-Rego et al. (1993)
43	Firca Galea	Lugo	43.51	-7.4	65			X						García-Amorena et al. (2008)
44	Montes del Bujo	Lugo	43.5	-7.4	500-600	¹⁴ C data at 7.8	X							Menéndez-Amor & Florschütz (1961), redrawn in López-García (1978)
45	Area Longa (Level I, II and III)	Lugo	43.60	-7.30	0	Oxygen isotope stages 5c, 4 and 3	X							Mary et al. (1977), Gómez-Orellana et al. (2007)
46	San Ciprián	Lugo	43.3	-7.05	802	Ages not provided	X							Muñoz-Sobrinho (2001)
47	Castro Chao de Samartín	Asturias	43.21	-6.88	675	Ages not provided	X							Ruiz-Zapata et al. (2003, 2005), redrawn in López-Merino (2009)
48	Pozo do Carballal	Lugo	42.70	-7.11	1330	c. >10.3 to present	X							Muñoz-Sobrinho et al. (1997)
49	Laguna de Lucenza	Lugo	42.58	-7.13	1375	c. 17.4 to present	X			X				Aíra-Rodríguez (1986), Santos-Fidalgo et al. (1997), Muñoz-Sobrinho et al. (2001)
50	A Golada	Lugo	42.70	-7.00	1225	¹⁴ C data at 3.4	X							Muñoz-Sobrinho et al. (1997)
51	Brañas de Lamela	Leon	42.76	-6.85	1280	¹⁴ C data at 3.1	X							Muñoz-Sobrinho et al. (1997)
52	A Cespedosa I and II	Lugo	42.87	-6.83	1415	¹⁴ C data at 2	X							Muñoz-Sobrinho et al. (1997)
53	Suarbol	Leon	42.86	-6.85	1080	¹⁴ C data at 1.25	X							Muñoz-Sobrinho et al. (1997)
54	Porto Ancares	Leon	42.86	-6.81	1580	Ages not provided	X							Muñoz-Sobrinho et al. (1997)
55	Pena Velosa	Asturias	43.01	-6.74	1350	c. 12 to present	X							Muñoz-Sobrinho et al. (2012)

Table 3. Continued.

Code	Site	Region	Lat.	Long.	Altitude (m a. s. l.)	Estimated chronology of palynological data (ka BP)	Palynology					Other palaeodata	References
							Plant macroremains	Sedimentary macrocharcoals	Sedimentary microcharcoals	Organic matter	Metals content/ Pb isotopes		
56	Puerto de Letiariegos	León	42.99	-6.41	1700	¹⁴ C dates at 11.8 and 8.9	X						García-Rovés et al. (2001), García-Rovés (2007), described in López-Merino (2009)
57	Villasaca de Laciama (Laguna del Castro)	León	42.95	-6.25	1317	c. 39.5 to 31 // 11.4 to 9 // ¹⁴ C data at >35 // 1 to present	X						Jalut et al. (2010)
58	La Mala (Laguna del Miro)	León	42.96	-6.21	1500	¹⁴ C data at 12.9 to present	X						Jalut et al. (2010)
59	Laguna de la Mata	Asturias	43.03	-6.15	1500	Ages not provided, probably c. 30 to present	X						Belet (1993), described in López-Merino (2009)
60	Lago de Ajo	Asturias	43.05	-6.15	1570	c. >14.3 to present	X	X					Allen et al. (1996)
61	Navia	Asturias	43.54	-6.72	10	No palynological data						X	García-Amorena et al. (2008)
62	Turbera Alto de la Espina	Asturias	43.38	-6.32	650	c. 10.7 to present	X				X		López-Merino (2009)
63	La Molina Mire	Asturias	43.38	-6.32	650	c. 2.7 to present	X						Cortizas et al. (2013, 2016)
64	Turbera de Las Dueñas	Asturias	43.56	-6.17	127	c. 9.9 to present	X					X	López-Merino et al. (2006), López-Dias et al. (2010, 2013a)
65	Turbera Huelga de Bayas	Asturias	43.57	-6.02	115	No palynological data				X		X	López-Dias et al. (2013a)
66	Turbera Monte Areo	Asturias	43.52	-5.76	200	c. >9.8 to present	X					X	López-Merino et al. (2010)
67	Estuario de Villaviciosa	Asturias	43.53	-5.39	15	c. >6.1 to present	X						García-Antón et al. (2006), García-Amorena et al. (2008)
68	Lago Enol	Asturias	43.27	-4.99	1070	c. 10.6 to 2.5	X				X		Menéndez-Amor (1975), López-Merino (2009), Moreno et al. (2010b, 2011)
69	Lago de la Ercina	Asturias	43.27	-4.98	1200	Ages not provided	X						Menéndez-Amor (1975)
70	Curueño	León	43.04	-5.41	1650	No palynological data		X					Fontbella-Blanco et al. (1999)
71	Turbera de Puerto de San Isidro	León	43.06	-5.38	1700	c. 8.3 to present	X						Carrión et al. (2012)
72	Cortegüero - S3	Asturias	43.11	-5.34	1530	¹⁴ C data at 5.7	X						Jiménez-Sánchez et al. (2003)
73	Brañagallones - S1	Asturias	43.12	-5.29	1230	¹⁴ C data at 29	X						Ruiz-Zapata et al. (2000, 2002b), redrawn in Carrión et al. (2012)
74	Mire in Pinar de Lillo	León	43.06	-5.25	1360	c. 1.7 to present	X						García-Antón et al. (1997), Muñoz-Sobrinó (2001)
75	Lillo II	León	43.06	-5.28	1500	c. >4 to present	X						Muñoz-Sobrinó (2001), Muñoz-Sobrinó et al. (2003)
76	Porma	León	43.05	-5.27	1300	No palynological data							Sánchez-Hernando et al. (1999)
77	Puerto de Tarna - S2	Asturias	43.09	-5.21	1415	¹⁴ C data at 20.6	X						Ruiz-Zapata et al. (2000), redrawn in López-Merino (2009), Jiménez-Sánchez et al. (2003)
78	Turbera de Comella / Comeya	Asturias	43.28	-4.98	834	¹⁴ C data at 40.5 // ¹⁴ C data at 8.3 // c. >3.8 to <3.1	X						Ruiz-Zapata et al. (2002c), redrawn in López-Merino (2009), Jiménez-Sánchez et al. (2003)
79	Esla	León	43.05	-4.83	1450	No palynological data							Sánchez-Hernando et al. (1999)
80	Turbera de Puertos de Riofrío	Cantabria	43.03	-4.69	1700	c. >10.2 to <2.2	X						Menéndez-Amor & Flörschütz (1963), redrawn in López-García (1978)

Table 3. Continued.

Code	Site	Region	Lat.	Long.	Altitude (m a.s.l.)	Estimated chronology of palynological data (ka BP)	Pollenology	Sedimentary macrocharcoals	Sedimentary microcharcoals	Organic matter	Metal content / Pb isotopes	Other paleodata	References
81	Turbera de Vidiago	Asturias	43.4	-4.66		Ages not provided	X						Menéndez-Amor (1950a, 1950b), redrawn in López-Merino (2009)
82	Turbera de Llano de la Mesa	Asturias	43.39	-4.63		Ages not provided	X						Menéndez-Amor (1950a, 1950b), redrawn in López-Merino (2009)
83	Turbera de Pendueles	Asturias	43.39	-4.63		Ages not provided	X				X		Menéndez-Amor (1950a, 1950b), redrawn in López-Merino (2009)
84	Turbera de Buelna	Asturias	43.39	-4.61	230	¹⁴ C dates at 2.2 and 1.7	X				X		Menéndez-Amor (1950a, 1950b), redrawn in López-Merino (2009); López-Días et al. (2013a)
85	Turbera de la Borbolla	Asturias	43.38	-4.61	50	No palynological data					X		Amor & Florschütz (1961), redrawn in López-Merino (2009); López-Días et al. (2013a)
86	Turbera de Llano Roñanzas	Asturias	43.38	-4.59	220	¹⁴ C dates at 3.2 and 1.4	X				X		García-Amorena et al. (2008); López-Días et al. (2013a)
87	Las Arenas - Tina Mayor	Cantabria	43.39	-4.51	0	¹⁴ C data at 10	X						Menéndez-Amor (1950a, 1950b); Mary & De Beaulieu (1973) redrawn in López-García (1978) and López-Merino (2009); Ortiz et al. (2008, 2010); Moreno et al., (2009); Gallego et al. (2013); López-Días et al. (2013a, 2013b)
88	Rio Bedema	Cantabria	43.41	-4.37	0	¹⁴ C data at 4.7	X						Mary et al. (1975)
89	Jerra I	Cantabria	43.39	-4.35	0	¹⁴ C data at 5.8	X						Mary et al. (1975)
90	Jerra II	Cantabria	43.40	-4.36	0	¹⁴ C data at 5.3	X						Mary et al. (1975)
91	Turbera Merón	Cantabria	43.38	-4.36	0	No palynological data							Mary (1990); García-Amorena et al. (2008)
92	Turbera de Oyambre	Cantabria	43.38	-4.33	0	No palynological data	X						Mary (1990); García-Amorena et al. (2008)
93	Turbera de Cullazón	Cantabria	43.23	-4.48	592	c. 3.6 to present	X						González-Pellajero et al. (2014)
94	Turbera del Pico Sertal	Cantabria	43.21	-4.43	940	c. 4.5 to present	X	X	X	X			Mariscal (1986); Carracedo et al. (2018)
95	Turbera de El Cueto de la Avellanosa	Cantabria	43.11	-4.36	1320	c. >6.1 to present	X	X	X	X			Mariscal (1983); Núñez (2018); Carracedo et al. (2016)
96	Turbera de La Molina	Cantabria	43.26	-3.97	484	c. 14.4 to present	X	X	X	X		X	Pérez-Obiol et al. (2016); Carracedo et al. (2016); Sánchez-Morales et al. (2022)
97	Turbera de El Cueto de la Espina	Cantabria	43.06	-3.91	1120	c. 5 to present	X	X	X	X			Rodríguez-Coterón (unpublished PhD); Salas (1993)
98	Turbera de Pico Año	Cantabria	43.1	-4.02	1288	4.1 to present	X						Mariscal (1993)
99	Turbera de Aisa	Cantabria	43.11	-4.01	560	c. 4.3 to 0.3	X						Iriarte-Chiapusso et al. (2003)
100	Santa Gadea	Burgos	42.9	-3.9	837	Ages not provided,	X						
101	Valle de La Nava I, II and III	Burgos	42.86	-3.89	870	probably c. 9–8.5 to present, c. >10 to <8.2	X						Menéndez-Amor (1988), redrawn in López-García (1978) and Muñoz-Sobrino (2001)
102	Puerto del Escudo	Cantabria	43.04	-3.87	940	¹⁴ C data at 7.8	X						Muñoz-Sobrino (2001)
103	Turbera de Estobac de Trueba	Cantabria	43.12	-3.70	1160	c. >5 to present	X						Mariscal (1987, 1989)
104	Turbera de Salamanca	Cantabria	43.24	-3.65	1290	c. 5.2 to present	X						Pérez-Díaz et al. (2016a)

Table 3. Continued.

Code	Site	Region	Lat.	Long.	Altitude (m a. s. l.)	Estimated chronology of palynological data (ka BP)	Palynology	Plant macroremains	Sedimentary microcharcoals	Sedimentary microcharcoals	Organic matter	Metal content / Pb isotopes	Other paleodata	References
105	Noja	Cantabria	43.46	-3.48	0	No palynological data	X							García-Amorena et al. (2008)
106	Tubera de Los Tomos I and II	Cantabria	43.15	-3.43	920	c. 7.8 to present	X						X	Peñalba (1994), Muñoz-Sobrimo (2001), Muñoz-Sobrimo et al. (2005)
107	Tubera de Zalama	Bizkaia	43.13	-3.40	1330	c. 7.2 to present	X				X			Pérez-Díaz et al. (2016b), Souto et al. (2016), Souto (2018)
108	Urdabai	Bizkaia	43.38	-2.68	5	c. 11.2 to >2.6	X							Iriarte-Chiapusso et al. (2006, 2016)
109	Tubera de Saldrogo I and II	Bizkaia	43.05	-2.71	625	c. 5.6 to present	X							García-Añón et al. (1989), Peñalba (1994)
110	Tubera de Belate	Navarra	43.03	-2.05	825	c. > 6.6 to present	X							Peñalba (1994), redrawn in Muñoz-Sobrimo (2001)
111	Estuary of Bidasoa	Gipuzkoa	43.3	-1.73	<200	¹⁴ C dates at 6.6 and 6.5	X						X	Sánchez-Gofí (1996)
112	Tubera de Añur	Navarra	43.25	-1.55	500	c. > 2.7 to present	X							Peñalba (1994), redrawn in Muñoz-Sobrimo (2001)
113	Lago de Arreo (ARRO4-1A-1K and ARRO4-2A-1K)	Araba	42.76	-2.98	655	c. 2.5 to present	X		X				X	Corella et al. (2013)
114	Ruidobro	Burgos	42.8	-3.6	835	¹⁴ C data at 0.45	X							Iriarte-Chiapusso et al. (2003)
115	Tubilla del Agua	Burgos	42.71	-3.80	765	No palynological data	X							Moreno-Amat & Iglesias-González (2009), García-Amorena et al. (2011)
116	Tubera de La Piedra	Burgos	42.63	-3.86	950	c. >12.3 to present	X							Ramírez-Rego et al. (1998)
117	Tubera de San Mamés de Abar	Burgos	42.68	-3.99	920	c. >10.7 to <3.7	X							Iriarte-Chiapusso et al. (2001, 2016), Muñoz-Sobrimo (2001)
118	Laguna Ciega	La Rioja	42.14	-2.68	1470	Ages not provided, probably c. 1.9 to present	X							Gil-García et al. (1995)
119	Hoyos de Iregua	La Rioja	42.02	-2.74	1780	c. 13.7 to present	X							Gil-García et al. (2002)
120	Laguna Masegosa	Soria	41.97	-2.82	1600	c. >9.5 to present	X				X			Von Engelbrechten (1998)
121	Laguna del Hornillo	Soria	41.97	-2.84	1620	c. 15.5 to present	X							Gómez-Lobo Rodríguez (1993), Gómez-Lobo Rodríguez et al. (1996), redrawn in Muñoz-Sobrimo (2001)
122	Laguna Negra	Soria	42.00	-2.84	1760	c. >9.6 to present	X				X			Von Engelbrechten (1998)
123	Lago Las Pardillas	Burgos	42.06	-3.00	1850	c. 9.3 to present	X				X			Sánchez-Gofí & Hannon (1999)
124	Nella hollow	Burgos	42.04	-3.01	1480	c. >5.2 to present	X				X		X	Von Engelbrechten (1998)
125	Quintanar de la Sierra	Burgos	42.03	-3.01	1470	c. >13.5 to present	X							Peñalba (1994), Peñalba et al. (1997)
126	Laguna grande (OS4)	Burgos	42.03	-3.05	1500	c. >17.15 to present	X							Ruiz-Zapata et al. (2002a)
127	Tubera Quintana Redonda	Soria	41.86	-2.84	1000	c. 9.3 to 6.7	X							García-Añón et al. (1994)
128	Tubilla del Lago	Burgos	41.80	-3.57	900	c. 6.5 to present	X			X				Moreno-Amat et al. (2009), García-Amorena et al. (2011), Morales-Molino et al. (2017)
129	Espinosa de Cerrato	Palencia	41.95	-3.93	885	c. 9.7 to present	X				X			Franco-Mugica et al. (2001), Morales-Molino et al. (2017)
130	El Carrizal	Segovia	41.32	-4.14	860	c. >8.7 to present	X		X					Franco-Mugica et al. (2005)

Table 3. Continued.

Code	Site	Region	Lat.	Long.	Altitude (m a. s. l.)	Estimated chronology of palynological data (ka BP)	Palynology	Plant macroremains	Sedimentary macrocharcoals	Sedimentary microcharcoals	Organic matter	Metal content / Pb isotopes	Other paleodata	References
131	Turbera de Brañuelas	León	42.6	-6.2	1000	¹⁴ C data at 8	X							Muñoz-Sobrinó (2001)
132	Xan de Llanas	León	42.30	-6.32	1500	c. >3.2 to present	X	X		X				Morales-Molino et al. (2011)
133	Arroyo de Vallefondo	León	42.28	-6.20	1000	c. >0.84 to present	X			X				Morales-Molino et al. (2011)
134	Arroyo de Ciervas	Zamora	41.96	-6.2	790	Ages not provided	X							Muñoz-Sobrinó (2001)
135	El Portillo I, II	Zamora	41.90	-6.13	870	Ages not provided	X							Muñoz-Sobrinó (2001)
136	Laguna de las Sanguijuelas ¹ / Sanabria Marsh ² / Lleguna ³	Zamora	42.11	-6.71	1050	c. >14 to present	X	X		X				Menéndez-Amor & Florschütz (1961) ¹ , Hannon (1986) ² , Turner & Hannon (1988) ² , Allen (1996) ² , Muñoz-Sobrinó (2001) ³
137	Laguna de las Sanguijuelas	Zamora	42.13	-6.70	1080	c. 14.8 to present	X							Muñoz-Sobrinó (2001), Muñoz-Sobrinó et al. (2004)
138	Laguna Cárdenas	Zamora	42.20	-6.75	1560	Ages not provided	X	X			X			Menéndez-Amor & Florschütz (1961)
139	Laguna Arroyas ¹ / de la Roya ²	Zamora	42.21	-6.78	1608	c. >12.9 to present	X							Menéndez-Amor & Florschütz (1961) ¹ , Allen et al. (1996) ² , Muñoz-Sobrinó et al. (2013, 2018) ²
140	Lagoa de As Lamas	Ourense	42.33	-7.33	1360	c. 12.8 to present	X							Maldonado-Ruiz (1994), Muñoz-Sobrinó (2001)
141	Fraga	Ourense	42.21	-7.27	1360	¹⁴ C data at 8	X			X				Santos-Fidalgo et al. (1997, 2000)
142	Serra de Queixa - H profile	Ourense	42.20	-7.24	1310	Ages not provided	X							Menéndez-Amor (1971), Sangos-Fidalgo (1982), Vidal-Romani & Santos-Fidalgo (1994)
143	El Maillo mire	Salamanca	40.54	-6.20	1100	c. 9.3 to present	X	X						Morales-Molino et al. (2013)
144	Lagoa Comprida 1	Portugal	40.36	-7.64	1600	¹⁴ C data at 9	X			X				Van der Knaap & Van Leeuwen (1997)
145	Lagoa Comprida 2	Portugal	40.36	-7.63	1645	Ages not provided	X							Van der Knaap & Van Leeuwen (1997)
146	Charco da Candieira	Portugal	40.34	-7.57	1409	c. >11.9 to present	X	X						Van der Knaap & Van Leeuwen (1995, 1997), redrawn in Muñoz-Sobrinó (2001)
147	Lagoa Claraça	Portugal	40.33	-7.60	1845	¹⁴ C data at 9.0	X							Van der Knaap & Van Leeuwen (1997)
148	Lagoa das Salgaideiras	Portugal	40.33	-7.60	1835	¹⁴ C data at 8.2	X							Van der Knaap & Van Leeuwen (1997)
149	Charca dos Cães	Portugal	40.33	-7.61	1795	¹⁴ C data at 6.4	X							Van der Knaap & Van Leeuwen (1997)
150	Chafariz do Rei	Portugal	40.33	-7.61	1770	Ages not provided	X							Van der Knaap & Van Leeuwen (1997)
151	Covaão do Boieiro	Portugal	40.33	-7.62	1730	¹⁴ C dates at 9.9 and 9.6	X							Van der Knaap & Van Leeuwen (1997)

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